

Handling clouds in assimilating high spectral resolution infrared radiances

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Content

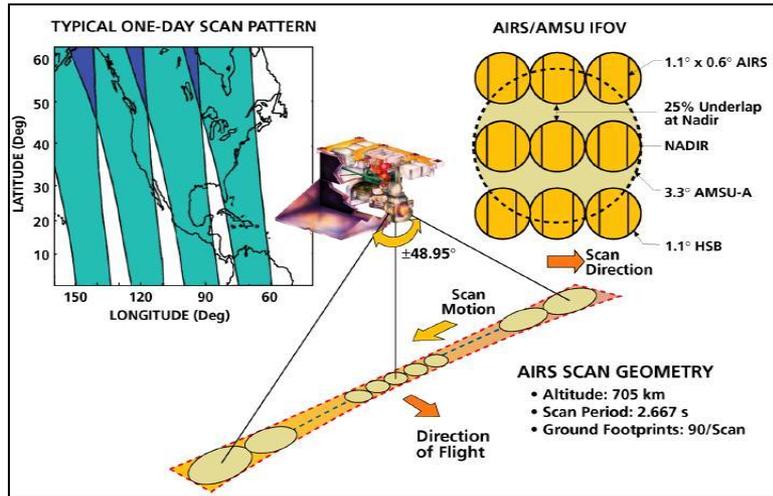
- Motivation
- Challenges on assimilating IR sounder radiances in cloudy situations
- Methodologies on handling clouds
- Assimilation experiments and findings
- Summary and future work

Motivation

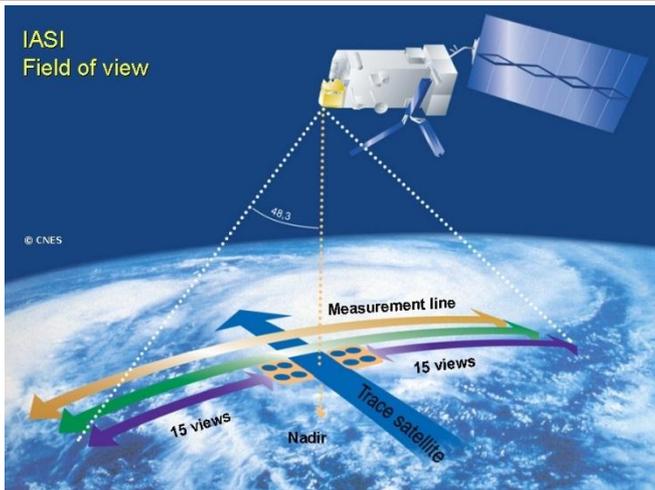
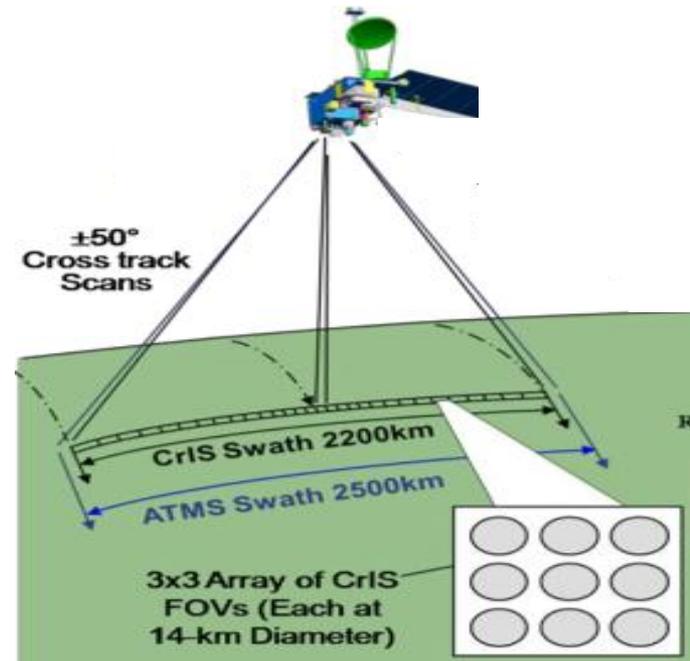
- Usually only clear IR channels (not affected by clouds) are used in most of data assimilation systems, and cloud contaminated channels have not been used effectively due to difficulties in modeling clouds in both forecast and radiative transfer models.
- Accurate cloud detection is very important for hyperspectral infrared (IR) radiance assimilation; a better cloud detection could reduce the data rejection and improve the assimilation.
- Since only 10 ~15% of hyperspectral IR footprints are clear globally, we need innovative ideas on assimilating radiances in cloudy situations.

Satellites with hyperspectral Infrared sensors

Aqua



Suomi NPP/JPSS

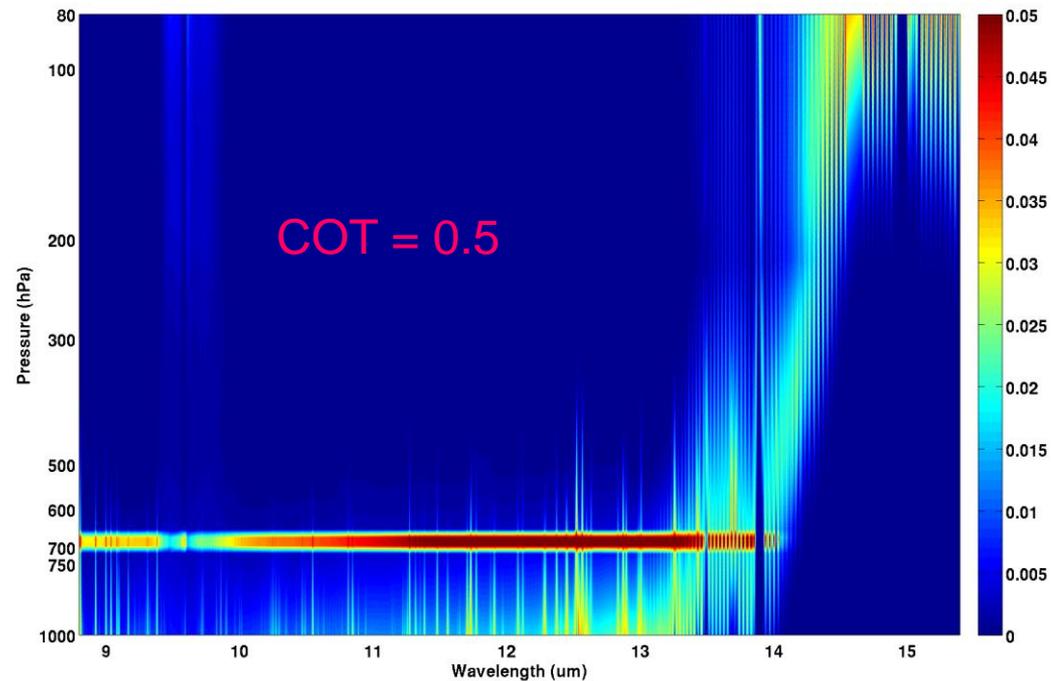
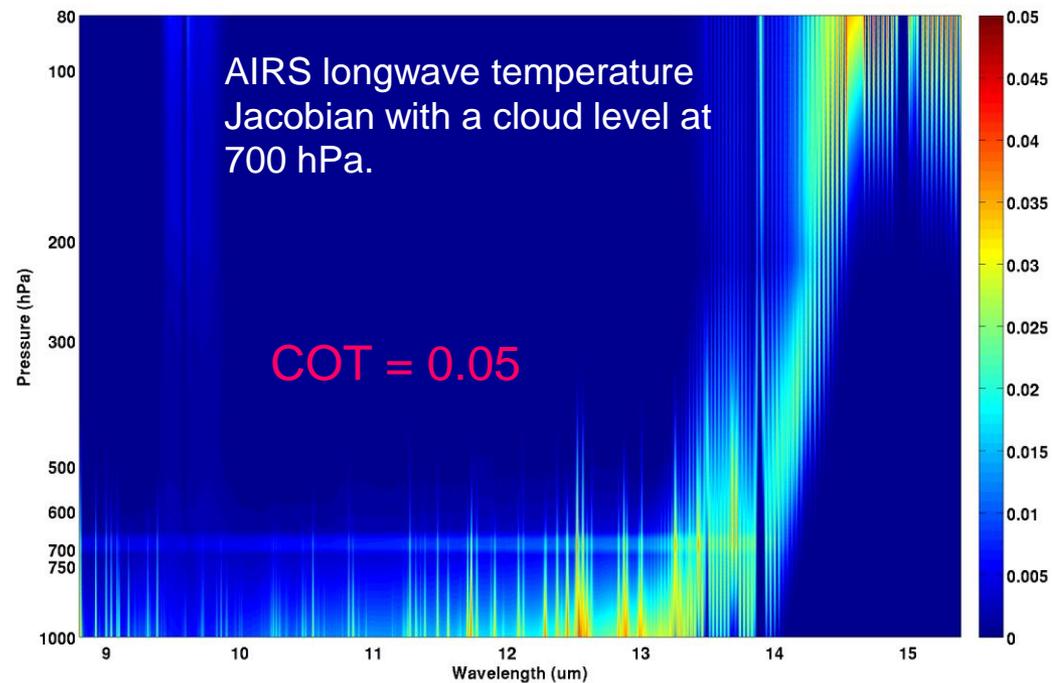


Metop

Challenges on assimilating radiances in cloudy situation:

(1) Both NWP and RTM have larger uncertainty;

(2) Big change of Jacobian at cloud level

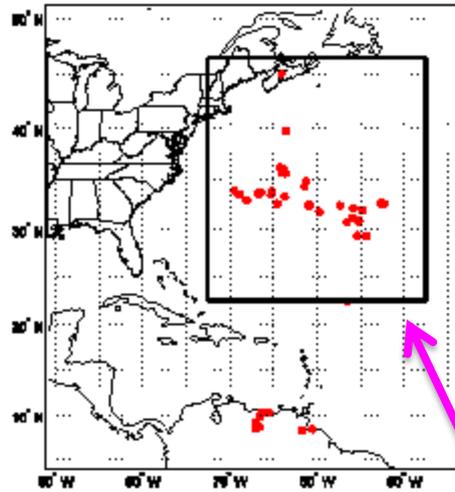
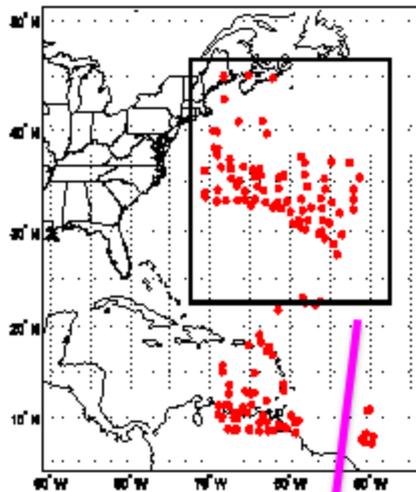


Handling clouds in IR radiance assimilation

- Black cloud assumption – rise the surface to cloud-top
- Hole hunting – excluding cloudy pixels
 - Stand-alone cloud detection with sounder data
 - Sub-pixel cloud detection with collocated high spatial resolution imager data
- Cloud-removal or cloud-clearing technique that transform the a cloud radiance spectrum to a clear equivalent radiance spectrum so that radiance assimilation can be treated as that in clear skies
- Direct assimilate radiance in cloudy skies – an accurate RTM in cloudy situation is very important

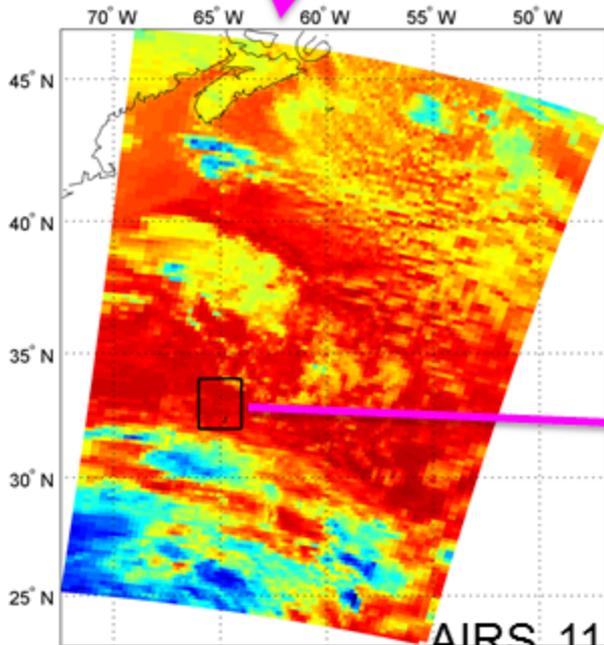
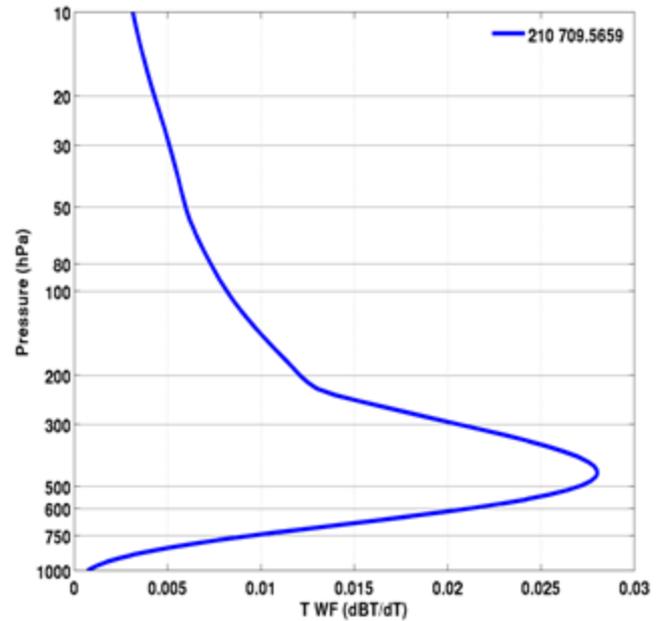
AIRS stand-alone cloud detection

MODIS cloud detection

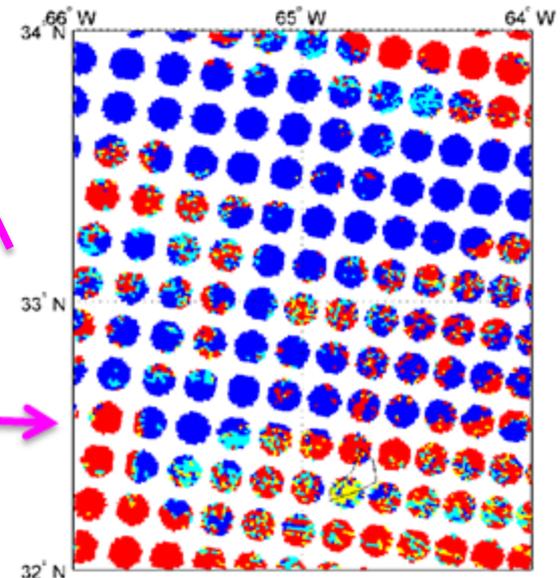


AIRS data at 06 UTC 25 October 2012 (Sandy)

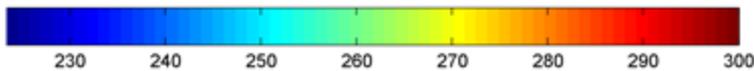
Channel Index 210, Wave number 709.5659



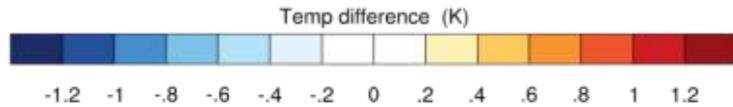
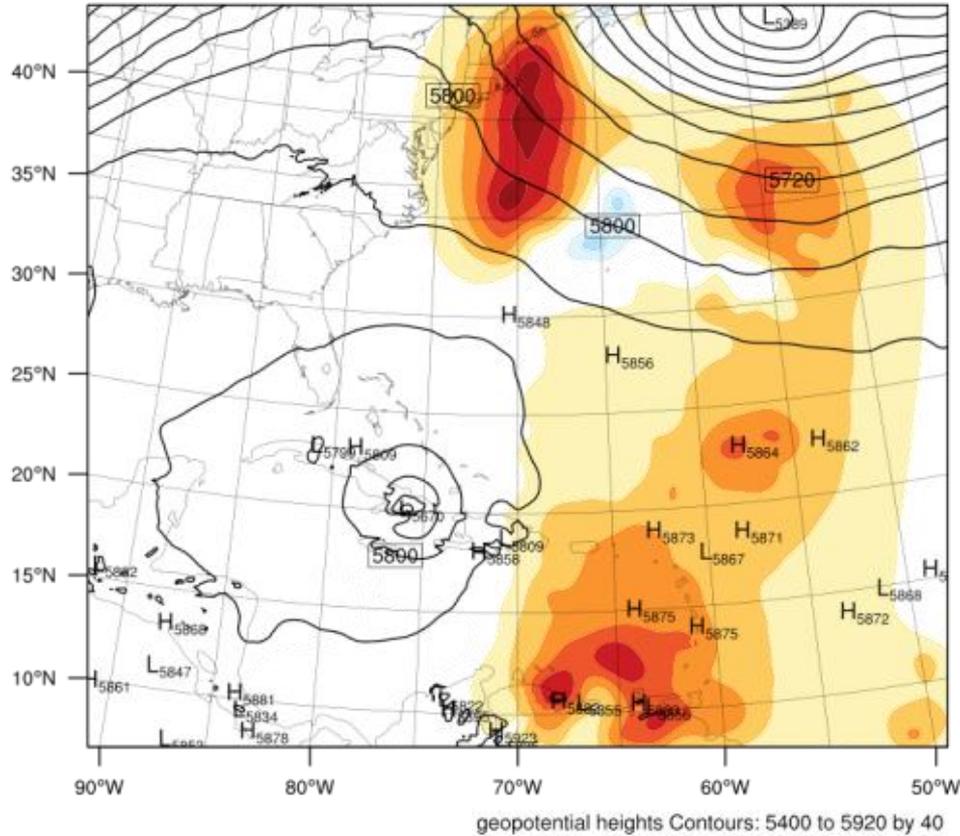
AIRS 11.3 μm BT (K)



AIRS sub-pixel cloud detection with MODIS



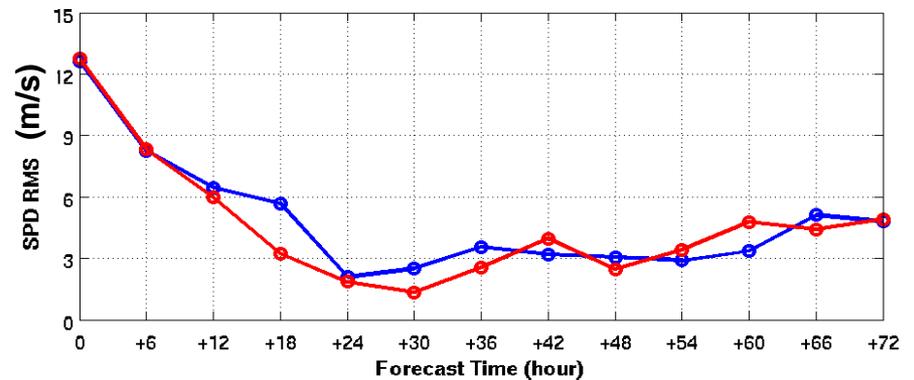
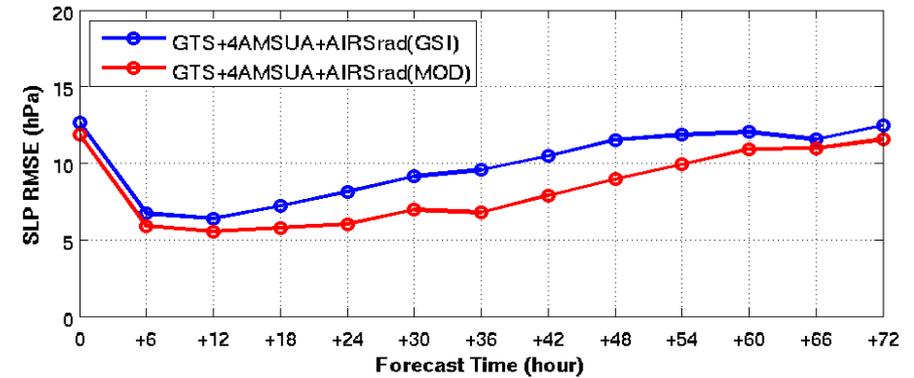
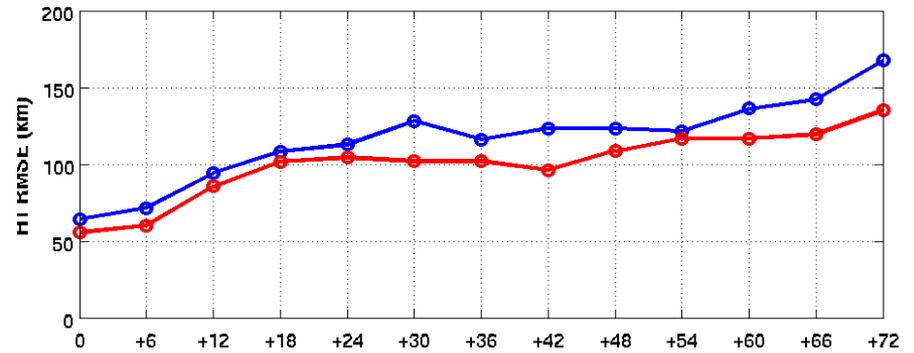
Temp difference (K)
geopotential heights (m) at 500 hPa



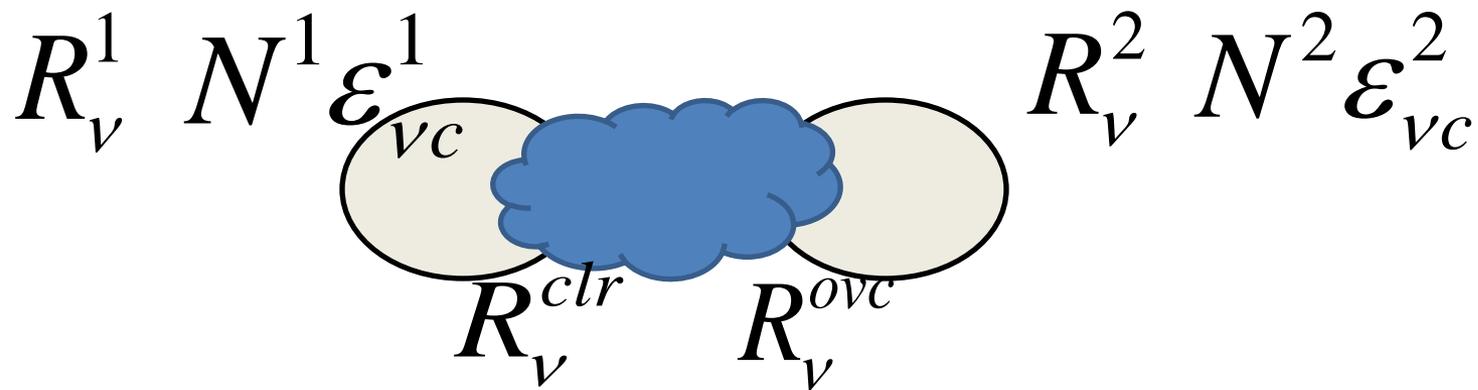
500 hPa temperature analysis difference
(AIRS(MOD) - AIRS(GSI))

Wang et al. 2014, GRL (accepted)

Hurricane Sandy (2012) forecast RMSE



72-hour forecasts of Sandy from 06z
28 to 00z 30 Oct, 2012



$$R_v^1 = (1 - N^1 \epsilon_{vc}^1) R_v^{clr} + N^1 \epsilon_{vc}^1 R_v^{ovc}$$

$$R_v^2 = (1 - N^2 \epsilon_{vc}^2) R_v^{clr} + N^2 \epsilon_{vc}^2 R_v^{ovc}$$

If we can estimate N^* (e.g. from MODIS) the cloud-cleared radiance R_v^{cc} can be obtained by

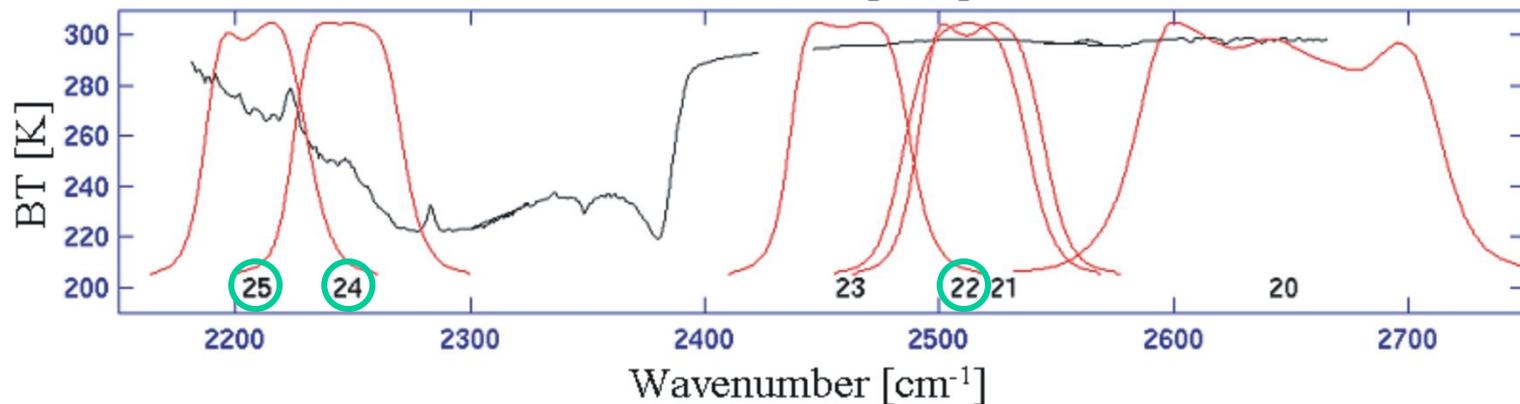
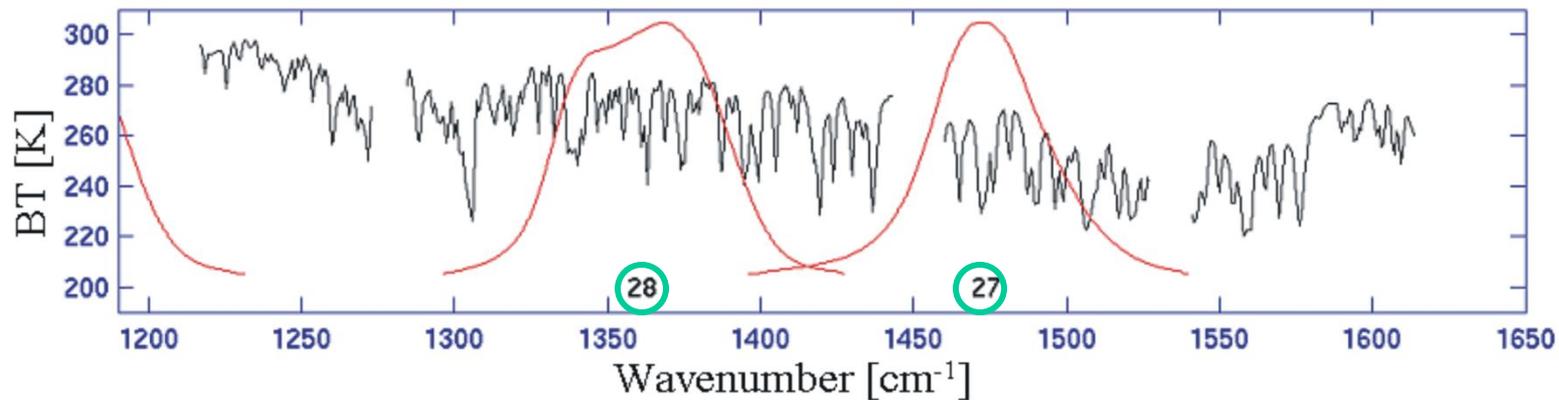
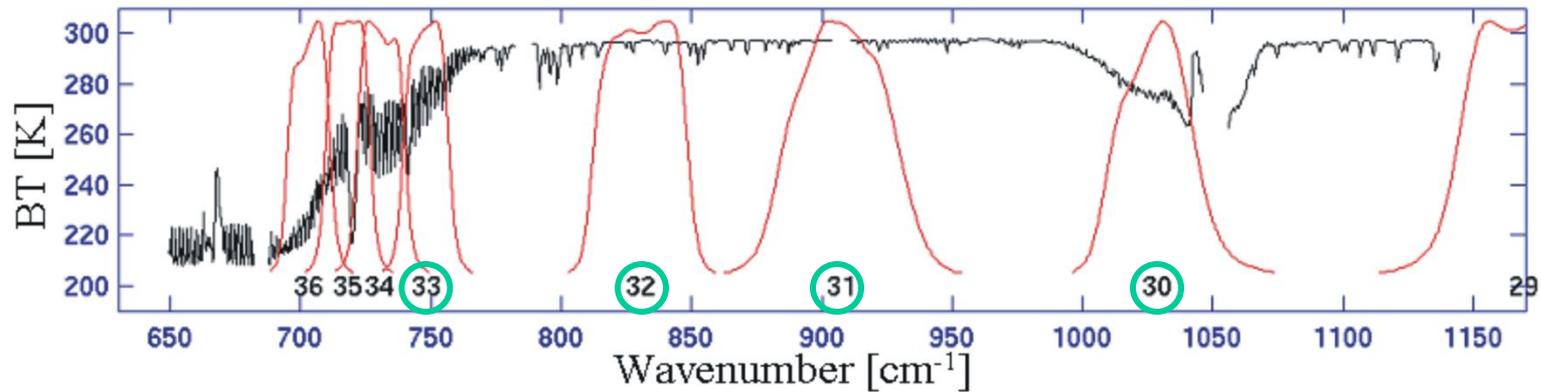
$$\rightarrow R_v^1 - R_v^{clr} = N^1 \epsilon_{vc}^1 (R_v^{ovc} - R_v^{clr})$$

$$R_v^2 - R_v^{clr} = N^2 \epsilon_{vc}^2 (R_v^{ovc} - R_v^{clr})$$

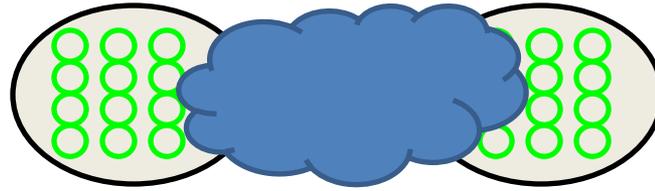
$$R_v^{cc} = \frac{R_v^1 - R_v^2 N^*}{1 - N^*}$$

$$\frac{R_v^1 - R_v^{clr}}{R_v^2 - R_v^{clr}} = \frac{N^1 \epsilon_{vc}^1}{N^2 \epsilon_{vc}^2} = N^*$$

Aqua MODIS IR SRF Overlay on AIRS Spectrum



Direct spectral relationship between IR MODIS and AIRS provides unique application of MODIS in AIRS cloud_clearing !

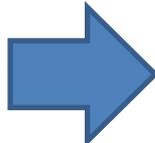


Using all MODIS IR bands (Li et al. 2005)

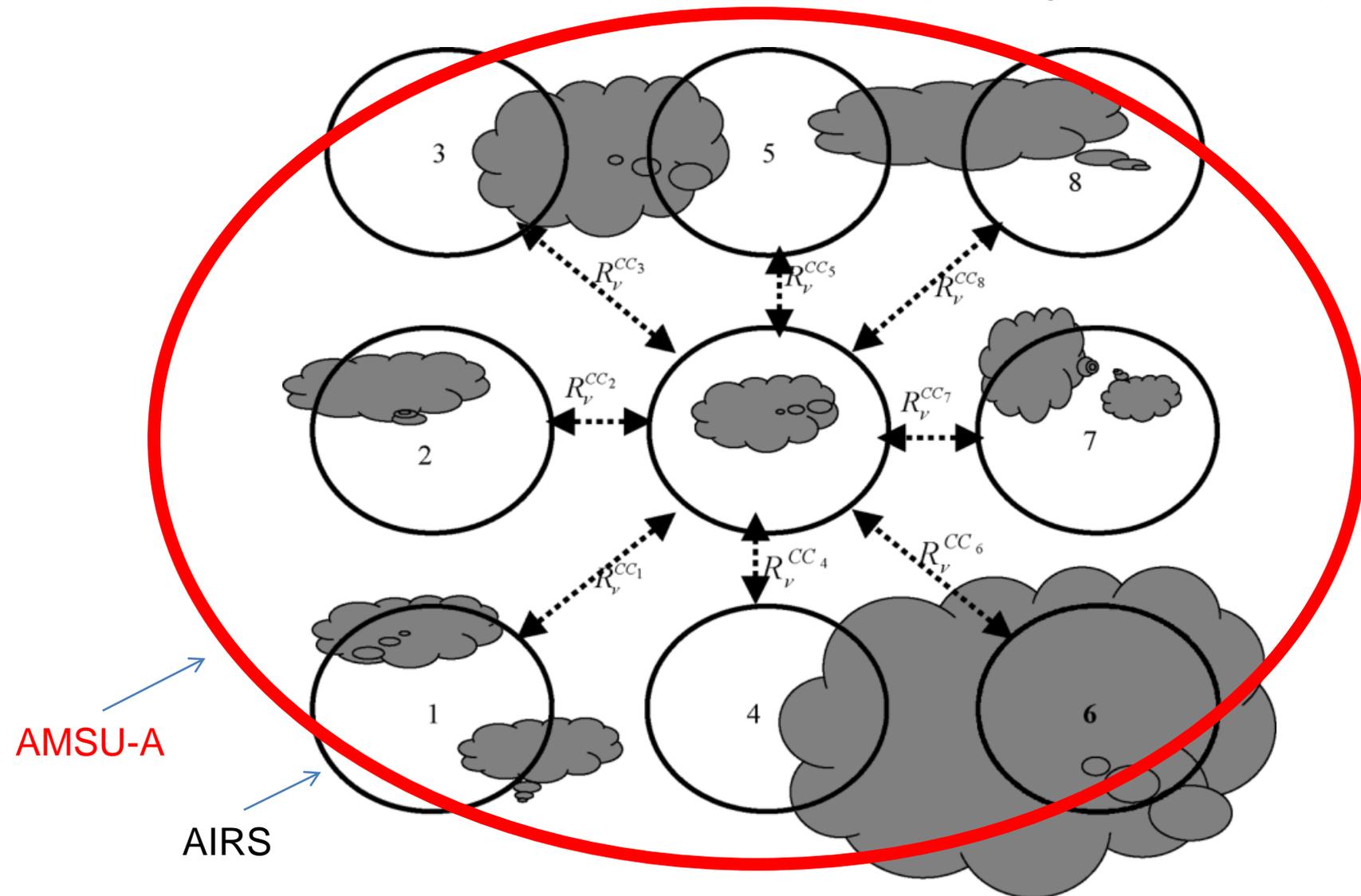
$$J(N^*) = \sum_i \frac{1}{\sigma_i^2} [(R_{M_i}^{clr} - f_i(R_v^{cc}))]^2 = \min$$

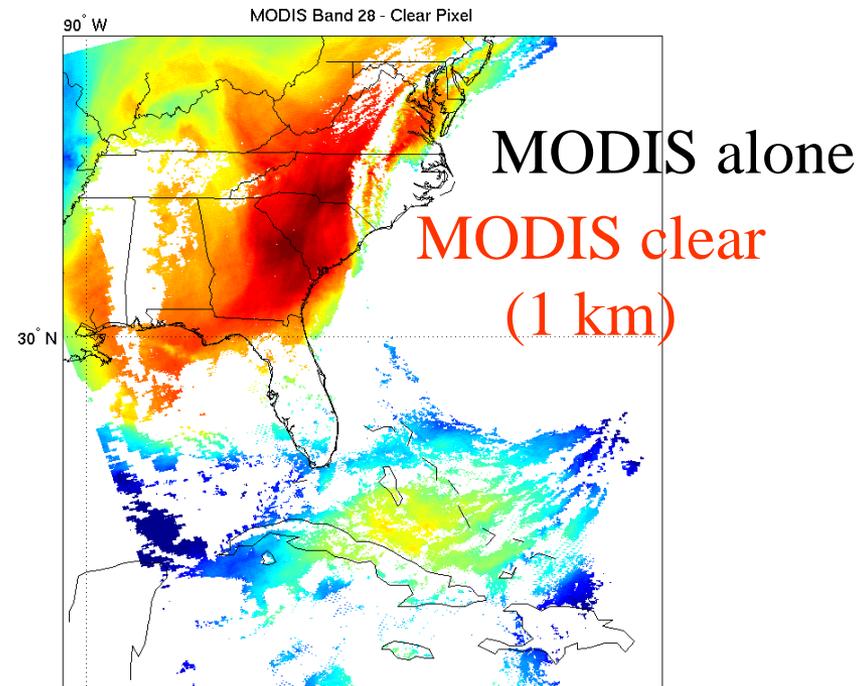
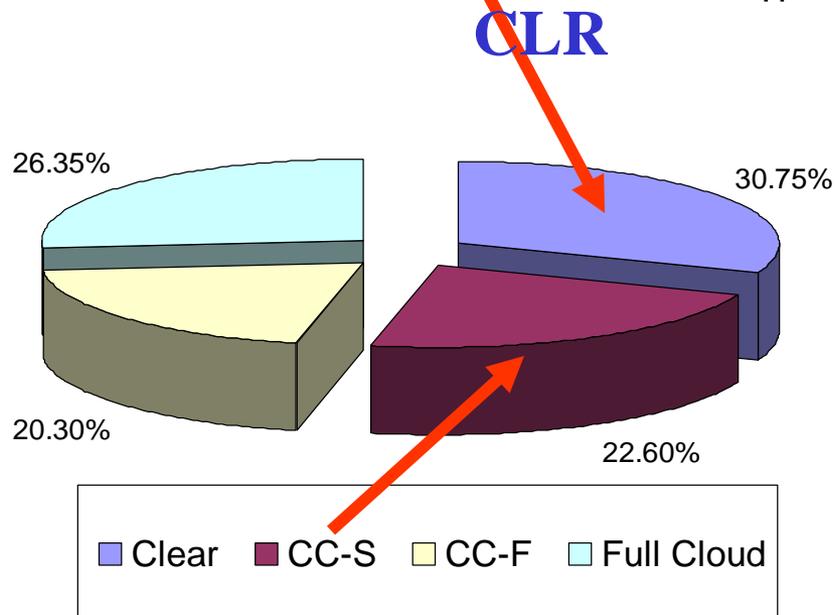
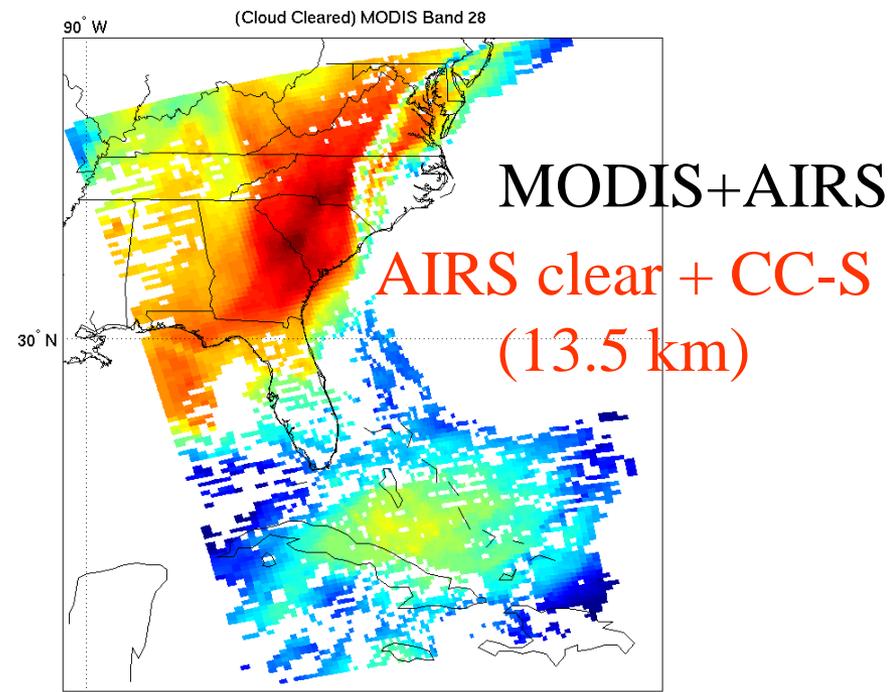
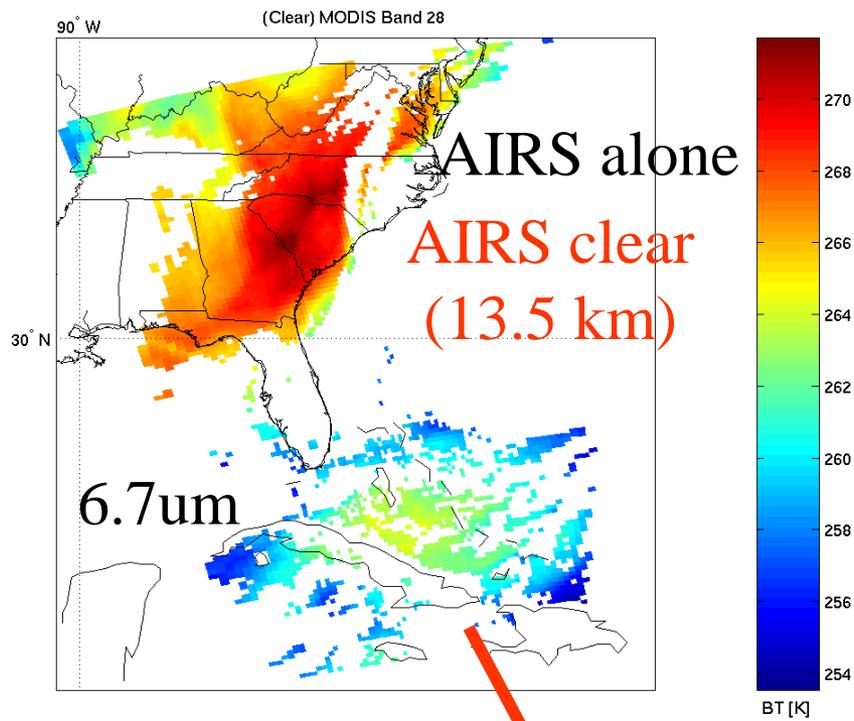
$$J(N^*) = \sum_i \frac{1}{\sigma_i^2} [(R_{M_i}^{clr} - f_i(\frac{R_v^1 - R_v^2 N^*}{1 - N^*}))] = \min$$

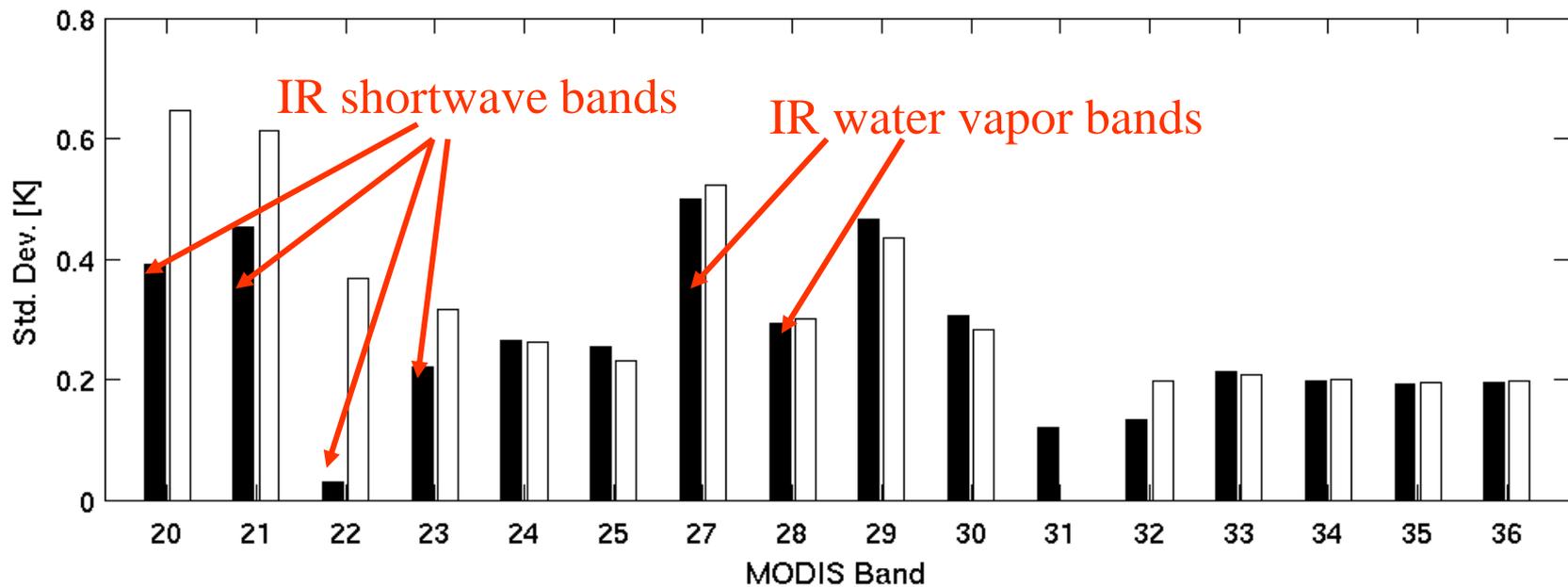
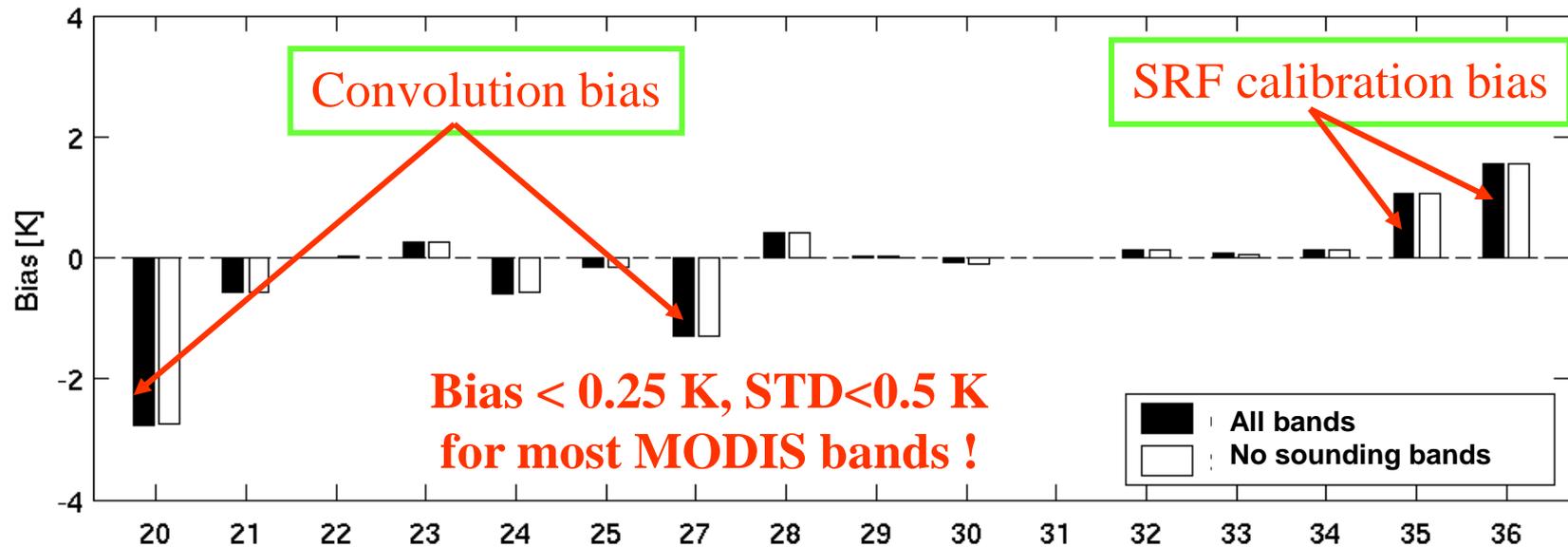
σ_i is NEdR for MODIS band i

Solve $\frac{\partial J(N^*)}{\partial N^*} = 0$  $N^* = \frac{\sum_i \frac{1}{\sigma_i^2} [f_i(R_v^1) - R_{M_i}^{clr}] [f_i(R_v^1) - f_i(R_v^2)]}{\sum_i \frac{1}{\sigma_i^2} [f_i(R_v^2) - R_{M_i}^{clr}] [f_i(R_v^1) - f_i(R_v^2)]}$

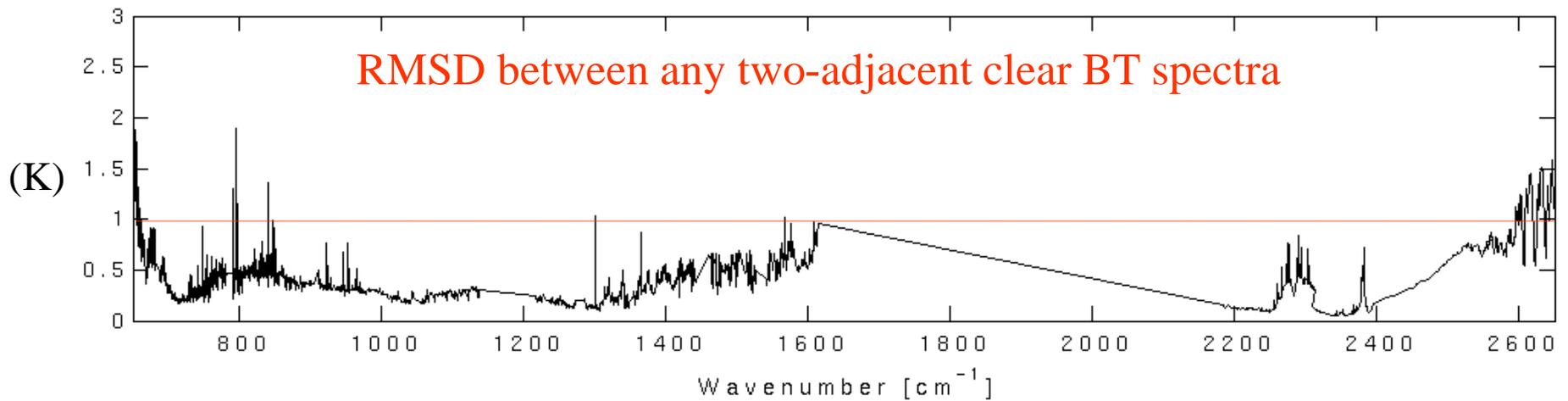
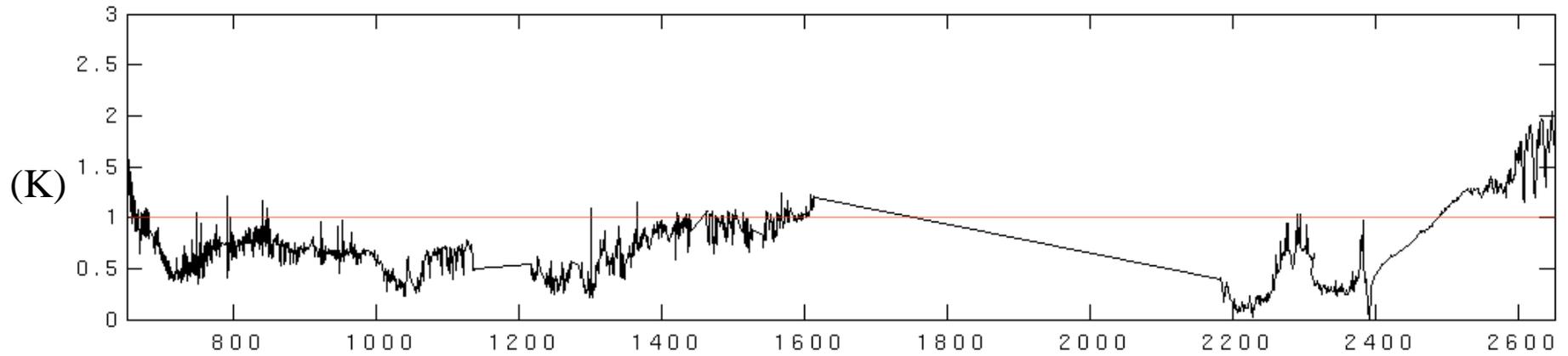
- (1) For each cloudy AIRS FOV, 8 pairs are used to derive 8 AIRS CC radiance spectra;
- (2) Compare AIRS CC radiances with MODIS clear radiance observations within the AIRS FOV, find the best pair and the corresponding CC radiance spectrum.



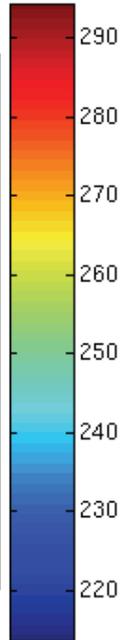
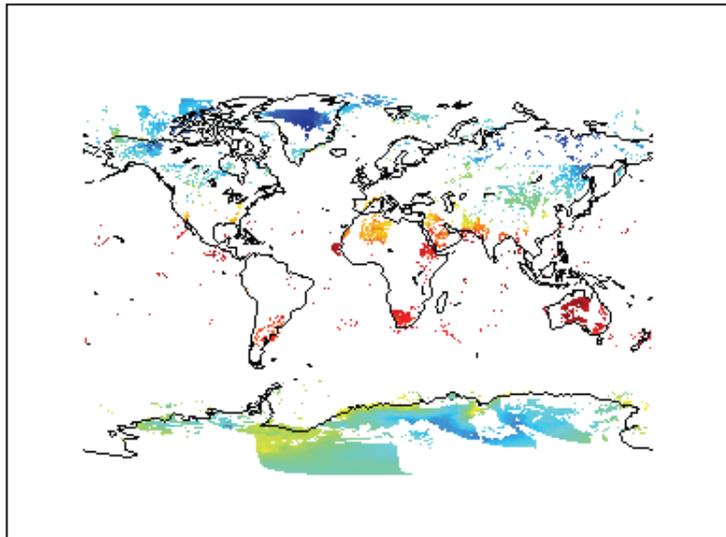




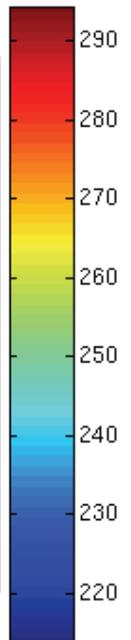
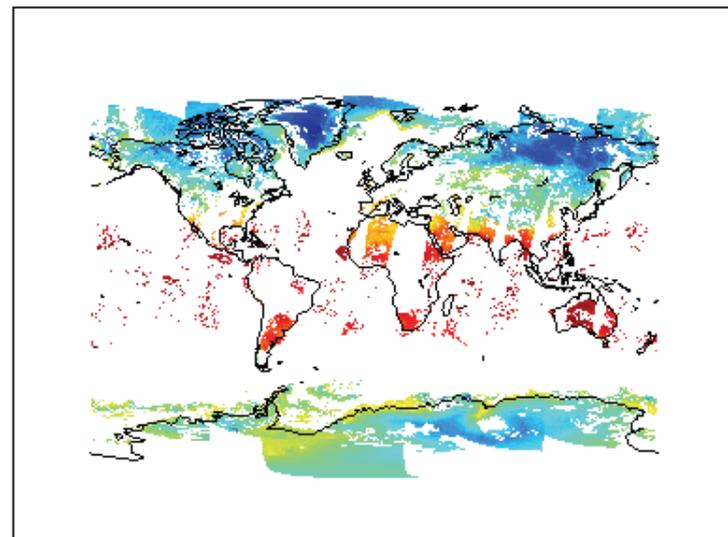
RMSD between CC BT spectra and their nearby clear BT spectra
(930 comparisons, AIRS granule 184 on 17 Sept. 2003)



Global AIRS clear (1000 cm⁻¹) Descending 20040101

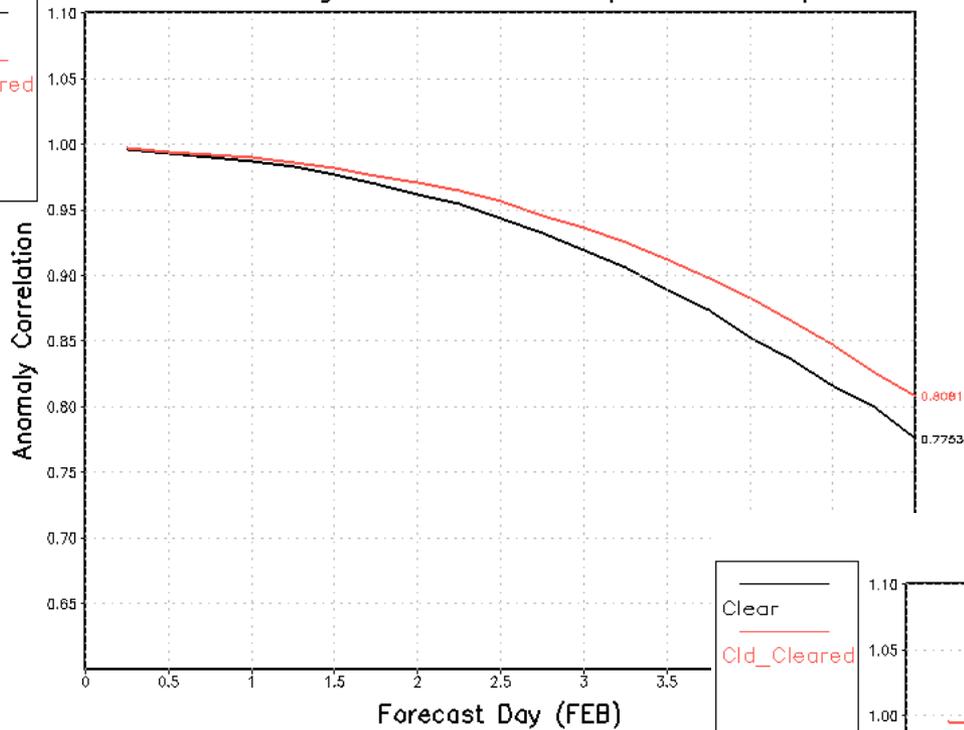


Global AIRS Clear & CC-S (1000 cm⁻¹) Descending 20040101



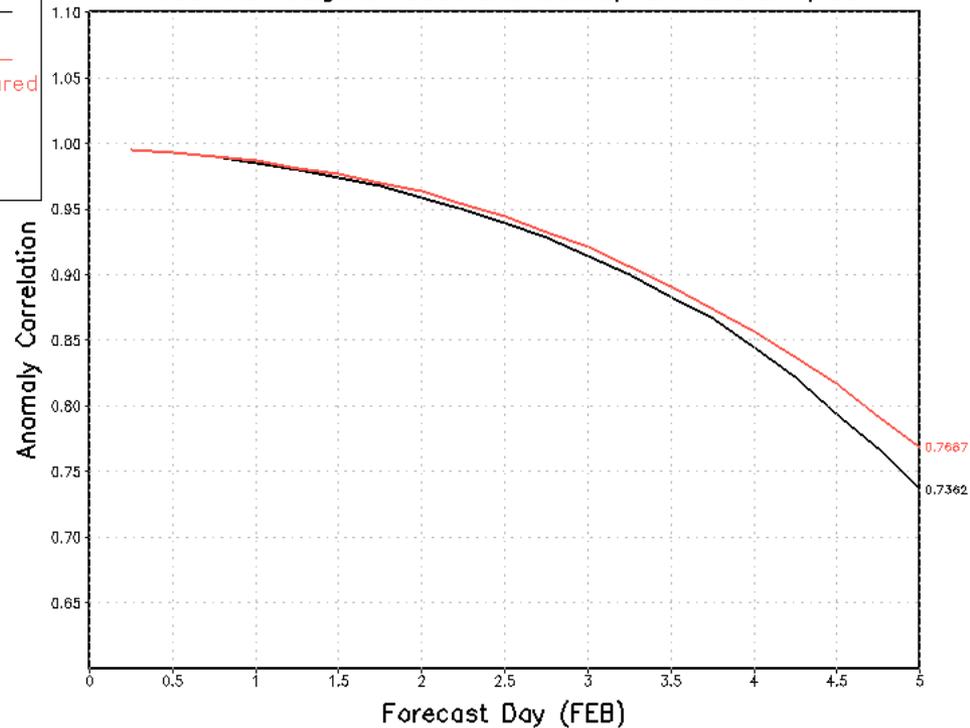
AIRS global clear and cloud clearing brightness temperature (descending) on Jan. 1, 2004.

500-mb Heights Northern Hemisphere ExtraTropics



Rienecker et al. 2008: GMAO's Atmospheric Data Assimilation Contributions to the JCSDA and future plans, *JCSDA Seminar*, 16 April 2008.

500-mb Heights Southern Hemisphere ExtraTropics



- GEOS-5 model resolution: $1^\circ \times 1.25^\circ \times 72L$
- Time frame: Jan 01 to Feb 15 2004
- Other Radiance data:
 - HIRS-2/HIRS3 (clear channels)
 - AMSU-A/EOS-AMSU-A
 - AMSU-B/MHS
 - SSM-I
 - GOES Sounders

Summary and future work

- Better cloud detection on AIRS with MODIS leads to significant NWP forecast improvement using GSI and WRF ARW systems;
- The approaches can be applied to CrIS/VIIRS, IASI/AVHRR;
- IR sounder cloud-clearing with collocated imager could expand the “clear” coverage for radiance assimilation;
- Future work will focus on assimilating the cloud-cleared IR sounder radiances in NWP.